

# CHEMICAL DEPOSITS ASSOCIATED TO METAVOLCANOSSEDIMENTARY SEQUENCES OF THE CENTRAL PORTION OF THE SÃO FRANCISCO CRATON IN THE STATE OF BAHIA, BRAZIL: A REVIEW

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**ABSTRACT** Supracrustal suites of metavolcanosedimentary rocks (SC) occur in the central portion of the São Francisco Craton in the State of Bahia. They were primarily formed by calc-alkaline to perialkaline volcanic flows and tuffs, clastic and chemical sediments. The latter are referred to the banded iron (BIF), iron-manganese (BIMF) and manganese (BMF) formations, iron-poor magnesian-rich carbonate formation (IPMRCF), and minor metacherts. BIF and IPMRCF are present in all the reported SC. BIF is typified by the presence of the oxide, carbonate and silicate facies. Occurrences of the sulfide facies are reported for the Sobradinho-Salitre and the Boquira SC. The burial metamorphism and thermal effects pronounced by the Transamazonian migmatization and granulitization are accountable for the generation of amphibolites, epidiosites, steatite and talc schists, phyllites, sandstones, quartzites, itabirites, metamorphic limestones to include banded carbonates, marbles and calc-silicate rocks. It is suggested that the deposition was similarly processed throughout an extensive crustal segment and that it was part of the cratonic evolution occurred during the end of the Archean to the beginning of the Proterozoic.

**Keywords:** São Francisco Craton., Paleoproterozoic, metavolcanosedimentary sequences, BIF, banded iron-manganese formations, magnesite

**INTRODUCTION** The Archean gneissic migmatitic terrains of Central Bahia are typified by the association with metavolcanosedimentary sequences which are in this paper termed supracrustal complexes and indicated by SC. These SC are named in accordance to their geographical distribution (Fig. 1). So, from the north to the south, they are called Sento Sé-Remanso, Casa Nova, Sobradinho, Salitre, Xique Xique, Boquira, Riacho de Santana, Caetité, Ibitira, Brumado, Umburanas, Contendas, Mirante, Guajerú, Urandí, Licínio de Almeida and Jacarací. They include a complex association of metavolcanic rocks and derived schists plus clastic and chemical sediments. The common geological features reported for the latter support the idea that they were generated by similar processes that took place from the end of the Archean to the beginning of the Proterozoic (Carvalho et al. 1997) and throughout an extensive but already individualized crustal segment. The structural relationship between the Archean basement and the SC suggests that they are associated to each other in a similar way to that one reported in the Rhodesian craton of Africa and the Yilgarn block of Australia for their substratum and the Archean greenstone belts (Phaup 1973, Turek and Compston 1970,

Mascarenhas 1976). During the Transamazonian a tectonothermal episode occurred in this portion of the S. Francisco Craton and it

was marked by migmatization and granulitization. A good evidence of these crustal recycling processes is presented by Pinto et al. (1998). The Middle Proterozoic Espinhaço Supergroup is deposited on these migmatized terranes of Central Bahia and it is stratigraphically succeeded by the Neoproterozoic São Francisco Supergroup. Younger unconsolidated Phanerozoic sediments cover the rest of the area.

With the purpose of shedding some more light on the knowledge of the chemical deposits, in this paper they are reviewed mainly

based on more recent data to including data from microanalyses of minerals. These analyses were performed at the Universidade de Campinas -UNICAMP by using an EDS equipment set to operate with 100 seconds exposure time, tension of 25 kV and I-probe of 7nA and at Universidade Federal da Bahia-UFBA by using a microprobe equipment CAMECA SX50.

**THE BASEMENT** The lithotypes of the cratonic basement include granulites, granoblastites, migmatites and gneisses that exhibit a wide range variation in composition being granodioritic and adamellitic the most common ones. Deposited on it are the SC. The intrusive plutons include granites, monzonites and syenites and they were emplaced from 2.2 to 1.7 Ga. (Jardim de Sá et al. 1976, Brito Neves et al. 1980, Mascarenhas and Garcia 1989). In more recent works, Paim et al. (1998) found an age of 2.1 Ga from Pb/Pb isochron obtained for the older syenitic lithofacies, termed Cara Suja, and Leahy et al. (1998) found the age of 1.7 Ga for the youngest magmatic pulses of the Cerâmica syenite. Both Cara Suja and Cerâmica syenites are part of the evolutionary history of the huge Guanambi-Urandi batholith that cuts and metasomatized the Urandi and Riacho de Santana SC. Pinto et al. (1998) found an age of 2.0 Ga from Pb-Pb isochron on monazite from a peraluminous granite which was generated from thermal reworking (block collision) of the Archean crust and intruded the Umburanas SC.

Remnants of the lithofacies of the SC that were deeply metasomatized and locally almost entirely eroded occur as lens like bodies that are spread all over the gneissic-migmatitic terranes. These metasomatized lithotypes are much more frequent at the proximities of the well preserved piles of the SC and near the metatexites and intrusive bodies.

## THE LITHOFACIES OF THE SC AND THE FACIES VARIATIONS OF THE CHEMICAL DEPOSITION

The SC comprise rocks derived from volcanic flows and tuffs, and from chemical and clastic sediments. The latter include grey to greenish phyllites, sandstones, quartzites and probably part of the schistose rocks. Locally, in the Sento Sé and Contendas-Mirante SC this suite of rocks predominates in the upper portion of the stratigraphic column (Souza et al. 1979, Marinho 1991). Mafic volcanic terms are reported at the lower units of the Contendas-Mirante (Marinho 1991) and Sobradinho (Silva and Cunha 1999) SC. Cunha and Froes (1994) reported spinifex textured komatiitic flows at the base of the



Figure 1 – General geology of the central portion of the S. Francisco Craton in the State of Bahia (modified after Inda and Barbosa 1978).

Umburanas SC. The intermediate to acidic volcanic rocks have an alkaline to perialkaline character. The volcanic rocks were almost entirely altered by the granitization-migmatization processes, into schistose materials, including quartz-mica, quartz-sericite, quartz-chlorite, quartz-biotite, muscovite, chlorite and garnet schists. Carvalho (1982) and Carvalho *et al.* (1982) described the rocks that were derived from the chemical sediments. From the shallow to the deep depositional zone these rocks include: metachert, ferruginous quartzites and orthoquartzites, itabirites, banded carbonates-amphiboles-martite and amphiboles-magnetite lithotypes. Table 1 summarizes some records for the chemical sedimentary facies.

Based on the chemical composition, the chemical metasediments are classified into the following groups of rock suites: banded iron formation (BIF), banded iron-manganese formation (BIMF), banded manganese formation (BMF), iron-poor magnesian-rich carbonate formation (IPMRCF) and metachert.

**The facies of BIF** The BIF are typified by the presence of the oxide, carbonate and silicate facies. The sulfide facies are reported for the Boquira SC where it constitutes de massive Pb-Zn ore (Carvalho *et al.* 1982,1997) and for the Sobradinho-Salitre SC (Silva and Cunha 1999). The oxide facies is composed of quartz-rich and iron oxides-rich bands which are alternately distributed within several scales of thickness (Carvalho and Albuquerque 1996). Hematite is the main iron oxide mineral and it may occur along with martite in proportions that tend to increase going towards the carbonate facies. Besides martite, the presence of amphiboles and iron-rich carbonates indicates this transitional zone. The presence of any other mineral is indicative of hydrothermal action. Figure 2 is the electronic micrograph obtained from a lithotype of the oxide facies that occurs at the Boquira mine nearby an apophysis of intrusive granite. The presence of biotite, apatite and ferrithorite as it is shown in this micrograph is an evidence of the hydrothermal effect on this facies. and in Figure 3 are the EDS spectra obtained for the minerals that are shown in Figure 2. The carbonate facies is basically composed of magnesian carbonate-rich bands that are altered to iron oxide (martite overwhelms hematite)-rich bands. Both kind of bands bear a varietal amount of amphiboles.

The metasomatism is much more pronounced in this facies than in the oxide one and this is due the more reactive character of the carbonates compared to both quartz and martite-hematite. The metasomatic effect may be observed by the generation of talc, amphiboles of the tremolite-actinolite series, diopside and scapolite. At the vicinities of metatexites and the intrusive granites, mineral changes may lead to the formation of calc-silicate lithotypes that represent true hornfels. Table 2 shows the composition obtained from electronic microprobe on carbonate and silicate minerals which are frequently reported in these and others studied lithofacies.

The FeO contents in the carbonate minerals rarely exceed 8.70 %, falling in the compositional field of Fe-dolomite. Figure 4 is a CaO-

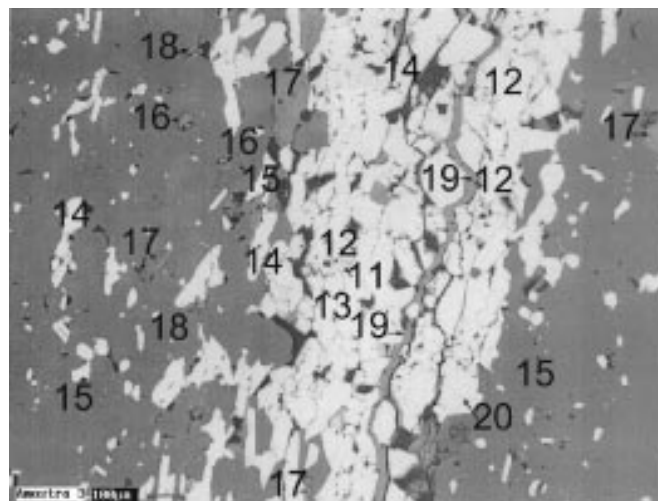


Figure 2 - EDS - Micrograph of oxide facies from the Boquira Complex. The analysed minerals are: Ferrithorite ( 11,13 ), Martite ( 12,20 ), Hematite ( 14 ), Quartz ( 15 ), Biotite ( 16,17 ), Apatite (18) and Calcite (19). Polished section.

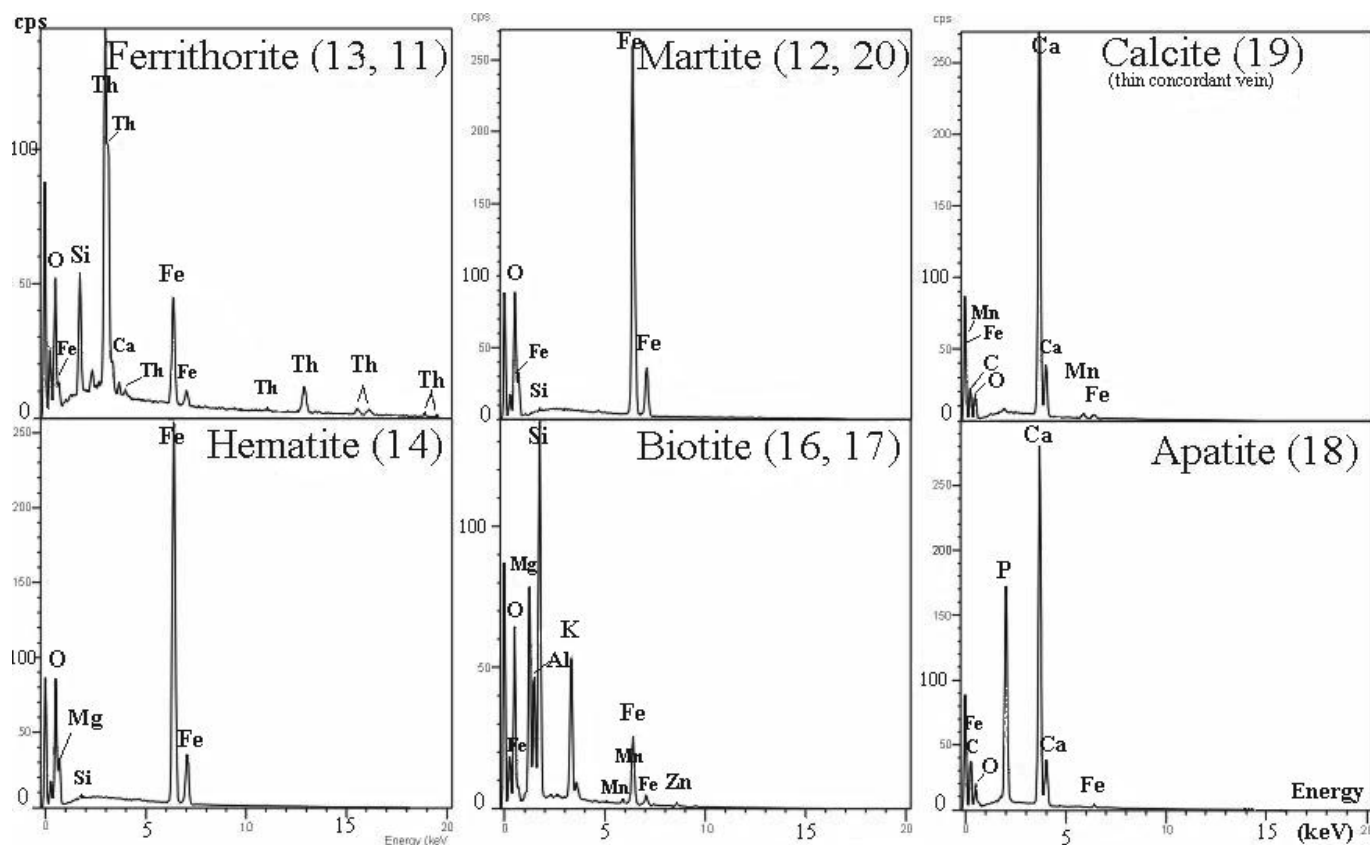


Figure 3 - EDS - Spectra from the minerals that are reported in the lithotype of the oxide facies of BIF shown in Figure 2.

Table 1- Some stratigraphic characteristics of the chemical lithofacies (CL) that occur in the central part of the S. Francisco Craton in the Central Bahia

HOST COMPLEX (SC)	HOST UNIT(S) OF THE CL	CHEMICAL NATURE AND MAIN LITHOTYPES	REFERENCES
SENTO SÉ-REMANSO-XIQUE XIQUE	Barreiro SC: <i>Baixa do Rancho Unit</i>	BIF. Metachert	Inda and Barbosa 1978 Souza et al.1979, Moraes Filho et al.1981,
	Colomé SC: <i>Serra do Choro Unit</i> <i>Castela Unit</i>	BIF- mostly oxide facies (itabirite). Metachert IPMRCF-thick banks of <i>magnetitic</i> and <i>dolomitic</i> marbles. Subordinate BIF and minor metacherts	
	<i>Serra da Capivara Unit</i>	BIF, IPMRCF, itabirite, metachert	
	<i>Barra Bonita Unit</i>	IPMRCF-dolomitic and calcitic marbles. Minor BIF	
CASA NOVA	<i>Macambira Unit</i>	Pyritic carbonates. Calc silicates and tremolites	Figueiroa and Silva F <sup>a</sup> 1990
	<i>Minadorzinho Unit</i>	BIF. Metachert	Angelim et al. 1993
	<i>Upper Unit</i>	IPMRCF - dolostones, marbles	Souza and Teixeira 1981
	<i>Lower Unit</i>	Pyritic metachert and IPMRCF -pyritized hornfels	Silva and Cunha 1999
SOBRADINHO-SALITRE	<i>Sobradinho Unit</i>	Pyritic and siliceous IPMRCF- pyritic carbonates. Silicate and sulfide facies. Metacherts. Hornfels. komatiitic mafic-ultramafic suites at the base	Angelim et al. 1993
BOQUIRA	Contendas Member	IPMRCF. Mainly dolomitic marbles	Kaul 1970, Carvalho et al.1982, Rocha 1983
	Boquira Member	BIF. All BIF's facies are present	
RIACHO DE SANTANA	<i>Upper Unit</i>	IPMRCF	Silva and Cunha 1999
	<i>Middle and Lower Units</i>	BIF. Komatiitic volcanic suite in the Lower Unit	
BRUMADO (only for the Serra das Éguas area)	Dolomitic Unit	Dolomite and lens of tremolite. Concordant lens like bodies of <i>magnetite</i>	Bondelos 1954
	Quartzitic Unit	Quartzites and BIF	
IBITIRA-BRUMADO	<i>Chemical Unit III</i>	BIF/itabirites. Minor banded metachert-carbonates	Oliveira et al. 1997
	<i>Chemical Unit II</i>	IPMRCF- Calcitic, dolomitic and magnetitic marbles	
	<i>Chemical Unit I</i>	BIF, IPMRCF. Interbedded mafic suites	
UMBURANAS	<i>Upper Unit</i>	IPMRCF	Cunha and Fróes 1994
	<i>Lower Unit</i>	BIF, IPMRCF- Metachert, calc silicate rocks Deposited on a komatiitic suite	
CONTENDAS-MIRANTE	<i>Upper Unit</i>	Scarce IPMRCF. Interbedded felsic volcanics	Marinho et al. 1979, 1980, Cunha et al. 1981
	<i>Lower Unit</i>	IPMRCF, BIF and metachert in association to fully differentiated volcanic flows and tuffs	
GUAJERÚ	<i>Middle and Upper Units</i>	BIF with predominance of the oxide facies (itabirite), and IPMRCF subordinate metadolomite	Bonfim et al. 1982
	<i>Lower Unit</i>	IPMRCF- marbles and metachert. Schistose komatiitic ultramafic suite at the base	
URANDÍ-CAETITÉ-CACULÉ-LICÍNIO DE ALMEIDA-JACARACÍ	<i>Quartzitic Unit</i>	BIMF, BIF, minor IPMRCF and Gondites	Ribeiro Filho 1973, Moraes et al. 1980, Machado 1983, Evangelista 1994
	<i>Schistose Unit</i>	BIMF, BMF, locally the presence of Mn-rich IPMRCF and lens-like bodies of BIF (itabirite and banded silicate facies)	

MgO-FeO diagram that shows the compositional distribution of the carbonates present in the BIF, BIMF and IPMRCF. The silicate facies is denoted by the alternations of amphiboles-rich with magnetite-rich bands. The more common types of amphiboles are those of the cummingtonite-grunerite, tremolite-actinolite and hornblende series. Amphiboles from the last two series may predominate in the places affected by the metasomatism caused by granitization and migmatization. Riebeckite was described by Souza (1992) in samples of this facies from the Sento Sé SC. The presence of amphiboles is not restricted to the silicate and carbonate facies. It may be locally reported in sub-banding of the oxide facies (Fig. 5.A).

**The facies of BIMF and BMF** The BIMF and BMF are only reported in the southwestern portion of this part of the São Francisco craton in the State of Bahia where they occur in the Urandi, Jacarací, Licínio de Almeida, Caculé and Caetité SC. They constitute the major important source of manganese ore in the State which is known as "The Mn-Ore District of Urandi-Licínio de Almeida-Caetité", that have been the subject of several detailed studies (e.g.: Ribeiro Filho 1973, Machado 1983, Basilio and Brondi 1986, Evangelista 1994). In this Mn-district both BIMF and BMF may be in association with BIF. This stratigraphic association is reported by the occurrence of individualized concordant stratum of BIF beneath the BMF and by interbedded BIF stratum in the BIMF deposit. Figure 5.B is the electronic micrography from a hematite-cryptomelane-rich lithotype of

the oxide facies of BIMF. The EDS spectra for the major minerals are shown in the Figure 6.

**The facies of IPMRCF** This kind of chemical sedimentary formation has a wide stratigraphical distribution. This is a striking characteristic of the SC which are considered in this paper, in spite of it may be limited to scarce narrow lens in some areas. This scarcity may be due to the "metasomatic assimilation" by the granitization and migmatization that took place in the older basement and that led to the formation of the calc-silicate hornfels.

**The siliceous facies** It occurs as metachert and quartzite and may contain accessory or varietal amounts of carbonate, iron oxide and silicate minerals that may be concentrated in poor defined bands. The different degree of recrystallization is indicative of the distinct metamorphic strains to which these metasediments were submitted. Some types of silicified rocks that are reported in association to the hornfels are the product of the siliceous metasomatism of the chemical sediments. This kind of metasomatism is also pointed out by Carvalho (1998) as being responsible for the talc generation in places where this process affect mafic-ultramafic rocks, the silicate and carbonate facies of BIF and BIMF and the carbonates of the IPMRCF.

**GEOCHEMICAL FACIES MODEL** As discussed by Carvalho and Albuquerque (1996), there is no major geochemical distinction

Table 2 - Chemical composition from electronic microprobe of some minerals which are reported in the lithofacies associated to the metavolcanosedimentary complexes of Central Bahia. The values are in percentage and they represent the average content of the element expressed in the oxide form. Each line corresponds to determinations on grain minerals found in the polished thin section made from the quoted lithofacies. nd = not detected.

COMPLEX	LITHOFACIES	MINERAL	MgO	Na <sub>2</sub> O	Al <sub>2</sub> O <sub>3</sub>	SiO <sub>2</sub>	FeO	NiO	BaO	K <sub>2</sub> O	CaO	Cr <sub>2</sub> O <sub>3</sub>	TiO <sub>2</sub>	MnO	Total
Sento Sé	BIF-Oxide F.	Hematite	nd	.06	.02	.60	89.17	.02	nd	.01	nd	nd	nd	.21	90.09
Sento Sé	BIF-Oxide F.	Hematite	nd	nd	.02	nd	89.78	nd	.05	.03	.03	nd	.01	.13	90.05
Boquira	BIF-Oxide F.	Hematite	nd	.06	.17	.34	89.36	.02	.08	.01	nd	.01	nd	nd	90.05
Boquira	BIF-Oxide F.	Hematite	nd	.01	.17	nd	89.78	.03	.05	nd	nd	nd	.05	.04	90.13
R.Santana	BIF-Oxide F.	Hematite	.05	.03	.61	nd	89.31	.03	nd	.02	nd	nd	.02	nd	90.07
R.Santana	BIF-Oxide F.	Hematite	.01	nd	.15	nd	89.78	.02	nd	.02	nd	.03	nd	nd	90.01
Ibit.-Brumado	BIF-Oxide F.	Hematite	.04	.13	.01	.05	89.56	nd	.14	nd	.02	nd	.06	.01	90.02
Caetité	BIMF-Oxide F.	Hematite	nd	nd	.01	nd	89.05	nd	nd	.01	nd	nd	nd	.69	89.76
Caetité	BIMF-Oxide F.	Magnetite	.03	.01	.09	.03	91.59	nd	nd	.04	nd	nd	nd	.73	92.52
Urandí	BIMF-Oxide F.	Magnetite	.02	.01	.14	.05	91.89	.02	nd	.05	nd	nd	nd	.65	92.83
Caetité	BIMF-Silicate F.	Magnetite	.06	nd	.34	.03	92.49	.02	nd	.04	nd	nd	nd	.65	93.63
L.de Almeida	BIMF-Silicate F.	Magnetite	.66	.06	.11	.19	91.49	.04	nd	.06	.05	nd	.08	.35	93.09
Sento Sé	BIF-Carbon. F.	Calcite	1.18	.01	nd	nd	1.21	nd	nd	nd	53.04	nd	.05	.87	56.36
Sento Sé	IPMRCF	Calcite	.76	.03	nd	.02	.70	.02	.02	.02	55.19	nd	.01	.85	57.62
Sento Sé	IPMRCF	Calcite	1.05	.04	.01	.02	1.04	.01	nd	.01	54.20	nd	.01	.97	57.36
Sento Sé	IPMRCF	Calcite	1.22	.05	nd	.01	.98	nd	.05	.01	54.64	nd	.01	.81	57.78
Sento Sé	IPMRCF	Calcite	1.26	.05	nd	.01	.99	.02	nd	.01	56.48	.01	nd	.61	59.44
Sento Sé	IPMRCF	Calcite	1.02	.03	.01	.04	.84	.01	.15	.02	53.87	nd	.01	.55	56.55
Sento Sé	IPMRCF	Calcite	1.02	.02	nd	.01	.91	.02	nd	nd	54.54	.02	nd	.65	57.19
R.de Santana	IPMRCF	Calcite	.89	.06	nd	nd	.83	nd	nd	nd	53.26	nd	nd	.58	55.62
R.de Santana	IPMRCF	Calcite	.93	.05	nd	.02	.84	.05	.04	.01	53.07	nd	nd	.58	55.55
Ibit.-Brumado	IPMRCF	Calcite	.98	.02	.01	.02	.85	.03	.02	.02	54.12	.01	nd	.66	56.74
Boquira	IPMRCF	Dolomite	22.15	nd	.03	.05	.55	nd	nd	nd	30.25	nd	nd	.02	53.05
Boquira	BIF-Carb. F.	Fe-Dolomite	16.32	.02	.04	.09	3.51	nd	.01	nd	28.72	nd	.02	.02	48.75
Caetité	BIF-Carb. F.	Fe-Dolomite	18.99	nd	.02	nd	8.26	.01	.01	nd	30.16	nd	nd	.88	58.33
Caetité	BIF-Carb. F.	Fe-Dolomite	18.72	.01	nd	nd	8.14	nd	.06	.01	30.39	.01	nd	.95	58.29
Caetité	BIF-Carb. F.	Fe-Dolomite	18.42	nd	.02	nd	8.14	nd	nd	nd	29.97	nd	nd	.96	57.51
Caetité	BIF-Carb. F.	Fe-Dolomite	18.52	nd	.01	nd	8.08	.03	nd	nd	30.60	nd	.01	1.04	58.29
Caetité	BIF-Carb. F.	Fe-Dolomite	20.58	.03	.03	.15	8.15	nd	nd	.01	28.19	.02	nd	.95	58.11
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	18.30	.01	nd	3.88	8.02	nd	nd	.01	30.35	nd	.01	1.06	61.64
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	18.25	.02	nd	3.63	8.27	.02	.10	nd	30.02	.02	nd	.10	60.43
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	17.32	.01	nd	nd	8.14	nd	nd	nd	30.03	.01	.01	.10	55.62
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	18.48	nd	.01	.02	8.34	nd	.03	nd	30.19	.01	nd	.98	58.06
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	19.18	.02	nd	nd	8.75	nd	nd	nd	30.18	.01	nd	.90	59.04
L.de Almeida	BIF-Carb. F.	Fe-Dolomite	18.97	.01	nd	.01	8.13	nd	.04	.01	30.12	.02	nd	1.06	58.37
Caetité	BIMF-Carb. F.	Fe-Ankerite	4.47	.04	.07	4.23	25.20	nd	nd	.02	11.27	.01	.01	.75	46.07
Caetité	BIMF-Carb. F.	Siderite	4.00	.04	.10	.11	32.04	nd	nd	.02	11.28	.01	.02	.82	48.44
Caetité	BIMF-Carb. F.	Siderite	3.38	.06	.16	.11	46.20	.01	nd	.01	10.56	nd	.02	.52	50.25
Sento Sé	IPMRCF	Magnesite	45.38	.01	nd	.32	nd	.02	nd	.01	.17	nd	.01	.15	46.07
Ibit.-Brumado	IPMRCF	Magnesite	46.31	.01	nd	.01	.41	nd	nd	nd	.35	nd	nd	.06	47.15
Ibit.-Brumado	IPMRCF	Magnesite	45.73	nd	nd	nd	.38	nd	.01	nd	.50	nd	nd	.10	46.72
Ibit.-Brumado	IPMRCF	Magnesite	44.96	nd	.01	.08	.32	nd	.01	nd	.91	nd	nd	.04	46.33
Ibit.-Brumado	IPMRCF	Dolomite	25.22	nd	.01	.02	.52	nd	nd	nd	30.40	nd	nd	.15	56.32
Sento Sé	BIF-Silicate F.	Trem./Actin.	18.00	.34	.51	57.18	8.01	nd	.04	.03	12.59	.02	.03	.24	96.99
Sento Sé	BIF-Silicate F.	Trem./Actin.	16.46	.81	2.64	55.17	8.97	.05	nd	.11	12.36	.01	.12	.28	96.98
Sento Sé	BIF-Silicate F.	Trem./Actin.	16.22	.90	3.12	55.00	8.80	nd	nd	.22	12.24	nd	.21	.25	96.96
Boquira	BIF-Silicate F.	Trem./Actin.	17.38	.61	1.47	56.85	7.69	nd	.10	.06	12.61	.02	.01	.21	97.01
Boquira	BIF-Silicate F.	Grunerite	.15	.32	.45	47.20	49.59	.01	nd	.05	.35	nd	.01	.28	98.41
Boquira	BIF-Silicate F.	Grunerite	.21	.30	.28	47.36	49.52	.01	nd	.04	.28	.01	.01	.30	98.32
Boquira	BIF-Silicate F.	Grunerite	.10	.21	.39	47.49	48.97	nd	.01	.02	.19	.01	.01	.17	97.57
R. de Santana	BIF-Silicate F.	Hornblende	13.80	.98	11.13	44.34	15.55	.03	nd	.32	11.14	.06	.80	.24	98.39
R. de Santana	BIF-Silicate F.	Hornblende	13.98	1.19	11.44	44.74	14.41	.03	nd	.37	11.13	.08	.83	.20	98.40
R. de Santana	BIF-Silicate F.	Hornblende	13.89	1.29	11.42	44.31	15.47	.03	nd	.32	10.62	.04	.79	.21	98.39

among the discussed BIF and either between them and those ones found in the Archean and Early Proterozoic terranes of Canada and South Africa. The BIFs here reported seem intermediate between the Archean Algoma and the Early Proterozoic Superior types. The previous discussions on REE distribution and on the similarities which are locally reported for the lithofacies pointed out four domains of distinct chemical deposition. However recent data as those records of the presence of pyritic facies and massive sulfides in the Sobradinho-Salitre SC and the evidence of the occurrence of this facies in the Riacho de Santana SC (Silva and Cunha 1999) make possible to

distinguish only two depositional areas of which one is marked by the absence of manganese deposition. The other depositional area is typified by the coeval iron-manganese depositions that are defined into the boundary of the Mn-District outlined by Basilio and Brondi (1986). In that area it is reported (a) a higher content of Mn in the minerals of the BIF's facies and (b) that the Mn content becomes higher in these minerals (hematite, magnetite, carbonates and silicates) if they occur in the facies that composes the BIMF. The presence of IPMRCF in the stratigraphic column of the SC seems to be one of the most significant identity of these terrains. Dolomite is the most

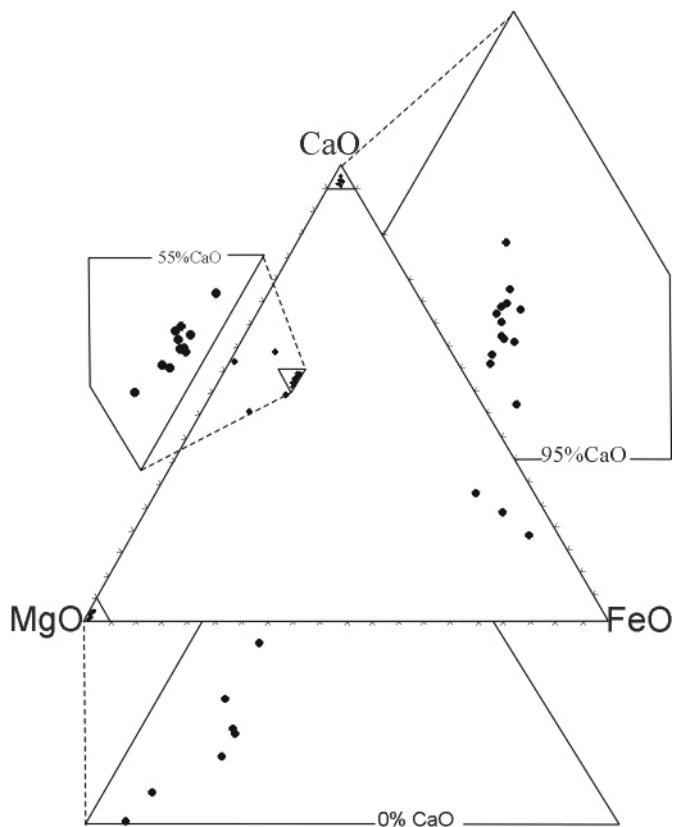


Fig. 4 - CaO-MgO-FeO diagram showing the distribution of the carbonate minerals from the carbonate facies of BIF, BIMF, and from IPMRCF (cf. Table 2).

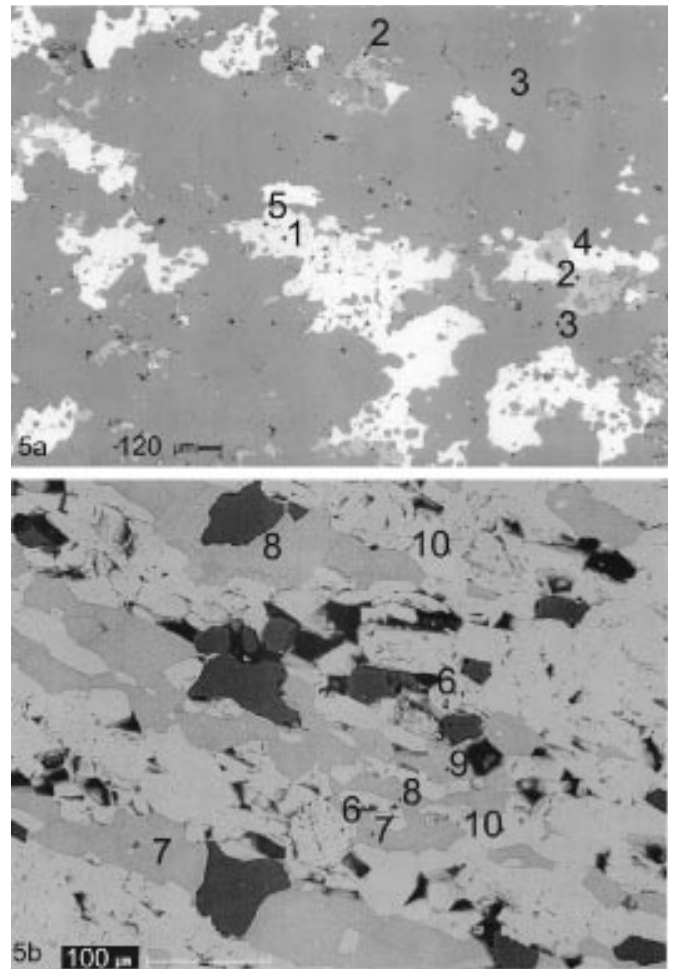


Figure 5 - EDS - Microphotographies from polished sections of (A): oxide facies from the Ibitira- Brumado Complex, showing an Amphibole-bearing sub-banding . The analysed minerals are: Hematite ( 1,4,5 ), Tremolite (2) and Quartz ( 3), and (B) of the oxide facies from the BIMF of the Caetité Complex. The analysed minerals are: Mn-Magnetite (6), Cryptomelane (7, 8, 9), Mn-Hematite ( 10) and Quartz ( dark grey ), Black = voids.

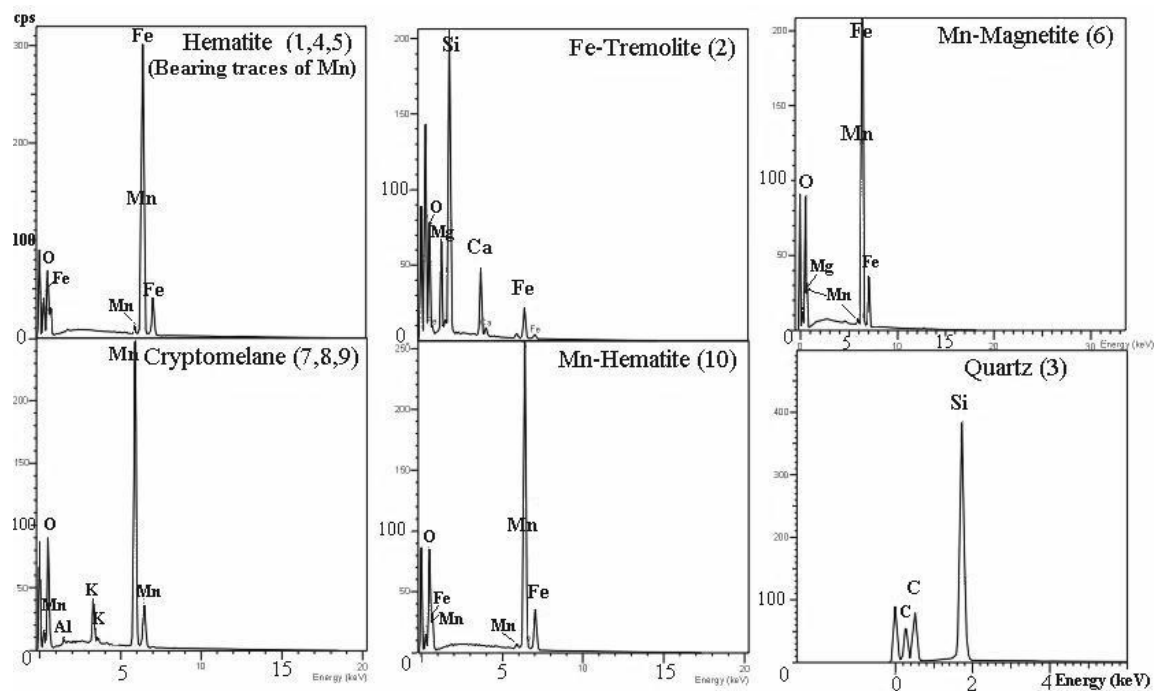


Figure 6 – EDS - Spectra for the minerals reported in the Figures 5.A (Hematite 1, 4, 5; Tremolite: 2, and Quartz: 3) and 5. B (Mn-magnetite: 6; Cryptomelane: 7, 8, 9 and Mn-Hematite: 10).

common carbonate mineral presents in the IPMRCF. The changes in the carbonate composition are locally noticeable by the facies variation as it is reported by the presence of stratiform levels of calcite or magnesite. The overall carbonate composition of the IPMRCF shows an iron content lower than that one reported for the minerals that compose the carbonate facies of BIF and BIMF which are Fe-dolomite, ankerite and siderite (Table 2 and Fig. 4).

The REE data presented by Carvalho and Albuquerque (1996) and Fonseca and Carvalho (1997) for the overall chemical metasediments yield  $(\text{Eu}/\text{Sm})_{\text{NC}}$ ,  $(\text{Sm}/\text{Yb})_{\text{NC}}$  and  $\Sigma\text{ETRL}/\Sigma\text{ETRP}$  values respectively of 0.47; 4.37, and 11.46. Also they show high

$\text{LREE}/\text{HREE}$  ratios and  $(\text{LREE})_{\text{NC}}/(\text{HREE})_{\text{NC}} > 1$ . That is an indication that the dissolved LREE were more efficiently coprecipitated by the original Fe-Mn oxide-hydroxide, carbonate, and silicate gels. The relative abundance of the IPMRCF is suggestive that  $\text{CO}_2$  played an important role on the pH control of the aqueous solution at the time of the deposition, when the  $\text{CO}_3^{2-}$  overwhelms the  $\text{HCO}_3^-$  species. This is a process that parallels the  $\text{P}_{\text{CO}_2}$  decreasing and the pH increasing and it also involves the increasing of the redox potential of the dissolved  $\text{Eu}^{++}$ , that rests in solution to explain the low  $(\text{Eu}/\text{Sm})_{\text{NC}}$  values found in the BIF and BIMF.

**CONCLUSIONS** Compared to the other chemical facies the BIF predominates in the Boquira SC and it is also quantitatively significant

in the Sento Sé-Remanso and Contendas-Mirante SC. The BIMF predominates in the Caculé and Licínio de Almeida SC. The BMF predominates in the Urandi SC and in many occurrences of the Caetité and Jacarací SC. The IPMRCF is the most abundant chemical lithofacies of the Sento Sé-Remanso and Ibitira-Brumado SC. The most frequent records of hornfels that are derived from BIF and IPMRCF are those reported for the Riacho de Santana SC. The geochemical signatures (data on stable isotopes, REE and other selected trace elements) define the BIF's type as intermediate from the Archean Algoma and the Paleoproterozoic Superior types (Carvalho *et al.* 1982, 1997). The depositional model may be alike to that one of plume applied to subaqueous region of active volcanism which is affected by intense leakage of gaseous-hydrothermal activities. The plume generation and chemical deposition took place throughout an extensive rifted zone of the craton and all the process may last from the end of the Archean to the Early Proterozoic. The local predominance of a given facies is an evidence that the environmental conditions were not homogeneous and not steadfast throughout such an extensive depositional zone.

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